

# Long Monthly European Temperature Series and the North Atlantic Oscillation

Jian Kang<sup>\*</sup>, Changli He<sup>\*\*</sup>, Annastiina Silvennoinen<sup>†</sup> and Timo Teräsvirta<sup>♠‡</sup>

<sup>\*</sup>Dongbei University of Finance and Economics, Dalian, P.R. China

<sup>\*\*</sup>Tianjin University of Finance and Economics, P.R. China

<sup>†</sup>NCER, Queensland University of Technology, Brisbane

<sup>♠</sup>Aarhus BSS, Aarhus University

<sup>‡</sup>C.A.S.E., Humboldt-Universität zu Berlin

Seminar at Eesti Pank

27 April 2023

# Introduction

Time series research on long temperature series (univariate)

- ▶ [Harvey and Mills \(2003\)](#) (the monthly Central England temperature (CET) series, 1723-2000; considerations based on time-aggregated quarterly and annual data)
- ▶ [Vogelsang and Franses \(2005\)](#) (the CET series, 1659-2000, and a Dutch (De Bilt) series, 1706-1993)
- ▶ [Proietti and Hillebrand \(2017\)](#) (the CET series, 1772-2013)
- ▶ [He et al. \(2019\)](#) (the CET series, 1772-2016)
- ▶ [Hillebrand and Proietti \(2017\)](#) (16 monthly European and 2 North American series, from the 18th century to (mostly) 2010s)

# Introduction

Time series research on long temperature series (multivariate)

- ▶ [He et al. \(2021\)](#) (3 monthly European and 2 Chinese series (Beijing and Shanghai), around 1830-2018)
- ▶ [He et al. \(in press\)](#) (20 monthly European series, around 1750-2015)

# Introduction

Perhaps the best known irregularly periodic climate phenomenon: the [El Niño Southern Oscillation](#) (ENSO) that has large effects on the weather in North and South America, and Australia.

- ▶ The strength of the ENSO is measured by the [Southern Oscillation Index](#) (SOI). The SOI is computed from fluctuations in the surface air pressure difference between Tahiti (in the Pacific) and Darwin, Australia (by the Indian Ocean)

# Introduction

European counterpart of ENSO: the **North Atlantic Oscillation** (NAO) that affects the weather in Europe.

- ▶ The NAO index is based on the surface sea-level air pressure difference between the Subtropical (Azores) High and the Subpolar (Iceland) Low.
- ▶ There are different definitions of the index, available for different time periods.

# Introduction

European counterpart of ENSO: the **North Atlantic Oscillation (NAO)** that affects the weather in Europe.

- ▶ The NAO index is based on the surface sea-level air pressure difference between the Subtropical (Azores) High and the Subpolar (Iceland) Low.
- ▶ There are different definitions of the index, available for different time periods.
- ▶ **In this work, we attempt at building a nonlinear time series model for characterising effects of the NAO on European temperatures over time (1823–2015).**

# The VSSMC-AR-X model

The model is an extended version of the Vector Seasonal Shifting Mean and Covariance Autoregressive (VSSMC-AR) model (He et al., in press), called VSSMC-AR-X model.

Notation:

- ▶  $\mathbf{y}_{Sk+s} = (y_{1,Sk+s}, \dots, y_{N,Sk+s})'$  is the  $N \times 1$  vector of endogenous variables
- ▶  $x_{Sk+s}$  the exogenous variable (can be generalised but only one exogenous variable in the application)
- ▶  $s = 1, \dots, S$  denotes the season (in the application the month)
- ▶  $k = 0, 1, \dots, K - 1$ , is the period (the year) counter
- ▶  $t = Sk + s$  is rescaled between zero and one, so the  $t$ th observation is indexed as  $u_{ks} = (Sk + s) / SK$ .

# The VSSMC-AR-X model

## Notation (continued)

- ▶  $D_{Sk+s}^{(j)}$  is a seasonal dummy variable:  $D_{Sk+s}^{(j)} = 1$  for  $j = s$ , zero otherwise.
- ▶ To maintain this notation for observations lagged by  $h$  seasons (months), i.e.,  $t = Sk + s - h$ , we adopt the modulo based equivalent representation  $t = S\tilde{k} + s_h$ .
- ▶ Thus,  $\tilde{k} = \lfloor (Sk + s - h - 1) / S \rfloor$  for  $k = 0, 1, \dots, K - 1$ , and  $s_h = s - h \pmod{S}$ . The residue system modulo  $S$  in this definition is the set  $\{1, \dots, S\}$ .



# The VSSMC-AR-X model

## The mean equation

The mean equation of the VSSMC-AR-X model is defined as follows:

$$\begin{aligned} \mathbf{y}_{S_{k+s}} &= \sum_{j=1}^S \{ \delta_j(u_{kj}) + \boldsymbol{\phi}_{j0} x_{S_{k+j}} \} D_{S_{k+s}}^{(j)} + \sum_{h=1}^p \{ \boldsymbol{\Phi}_h \mathbf{y}_{S_{\tilde{k}+s_h}} \\ &\quad + \sum_{j=1}^S \boldsymbol{\phi}_{jh} x_{S_{\tilde{k}+j_h}} D_{S_{k+s}}^{(j)} \} + \boldsymbol{\varepsilon}_{S_{k+s}} \\ &= \delta_s(u_{ks}) + \boldsymbol{\phi}_{s0} x_{S_{k+s}} + \sum_{h=1}^p \{ \boldsymbol{\Phi}_h \mathbf{y}_{S_{\tilde{k}+s_h}} + \boldsymbol{\phi}_{sh} x_{S_{\tilde{k}+s_h}} \} \\ &\quad + \boldsymbol{\varepsilon}_{S_{k+s}}. \end{aligned} \tag{1}$$

# The VSSMC-AR-X model

## The mean equation

In (1),

- ▶  $\Phi_h$  is an  $N \times N$  parameter matrix, and
- ▶  $\phi_{jh}$ ,  $h = 0, 1, \dots, p$ ;  $j = 1, \dots, S$ , are  $N \times 1$  parameter vectors.
- ▶  $\varepsilon_{Sk+s}$  is the  $N \times 1$  vector of independent errors with  $E\varepsilon_{Sk+s} = \mathbf{0}$  and  $E\varepsilon_{Sk+s}x_{Sk+s} = \mathbf{0}$ .
- ▶  $x_{Sk+s}$  is at least weakly exogenous to the parameters in (1).

# The VSSMC-AR-X model

## The mean equation

In (1),

- ▶  $\Phi_h$  is an  $N \times N$  parameter matrix, and
- ▶  $\phi_{jh}$ ,  $h = 0, 1, \dots, p$ ;  $j = 1, \dots, S$ , are  $N \times 1$  parameter vectors.
- ▶  $\varepsilon_{Sk+s}$  is the  $N \times 1$  vector of independent errors with  $E\varepsilon_{Sk+s} = \mathbf{0}$  and  $E\varepsilon_{Sk+s}x_{Sk+s} = \mathbf{0}$ .
- ▶  $x_{Sk+s}$  is at least weakly exogenous to the parameters in (1).
- ▶ Roots of  $|\mathbf{I}_N - \sum_{h=1}^p \Phi_h z^h| = 0$  lie outside the unit circle.

# The VSSMC-AR-X model

## The mean equation

The deterministic time-varying intercept vector of the VSSMC-AR-X model for season  $s$  equals

$\delta_s(\mathbf{u}_{ks}) = (\delta_{s1}(\mathbf{u}_{ks}), \dots, \delta_{sN}(\mathbf{u}_{ks}))'$ , where the  $s$ th time-varying coefficient  $\delta_{sl}(\mathbf{u}_{ks})$  of equation  $l$  is defined as

$$\delta_{sl}(\mathbf{u}_{ks}) = \delta_{sl0} + \sum_{i=1}^{q_{sl}} \delta_{sli} \mathbf{g}_{sli}(\mathbf{u}_{ks}; \gamma_{sli}, c_{sli}). \quad (2)$$

# The VSSMC-AR-X model

## The mean equation

The deterministic time-varying intercept vector of the VSSMC-AR-X model for season  $s$  equals

$\delta_s(u_{ks}) = (\delta_{s1}(u_{ks}), \dots, \delta_{sN}(u_{ks}))'$ , where the  $s$ th time-varying coefficient  $\delta_{sl}(u_{ks})$  of equation  $l$  is defined as

$$\delta_{sl}(u_{ks}) = \delta_{sl0} + \sum_{i=1}^{q_{sl}} \delta_{sli} g_{sli}(u_{ks}; \gamma_{sli}, c_{sli}). \quad (2)$$

In (2), the transition function is either logistic,

$$g_{sli}(u_{ks}; \gamma_{sli}, c_{sli}) = (1 + \exp\{-\gamma_{sli}(u_{ks} - c_{sli})\})^{-1} \quad (3)$$

or exponential,

$$g_{sli}(u_{ks}; \gamma_{sli}, c_{sli}) = 1 - \exp\{-\gamma_{sli}(u_{ks} - c_{sli})^2\} \quad (4)$$

In (3) and (4),  $\gamma_{sli} > 0$ , for  $i = 1, \dots, q_{sl}$  and  $s = 1, \dots, S$ .

# The VSSMC-AR-X model

## The covariance equation

The error term  $\varepsilon_{S_{k+s}}$  of the VSSMC-AR-X model is decomposed as  $\varepsilon_{S_{k+s}} = \Sigma_{S_{k+s}}^{1/2} \zeta_{S_{k+s}}$ , where  $\zeta_{S_{k+s}} \sim \text{iid}(\mathbf{0}, \mathbf{I}_N)$ , and

$$\Sigma_{S_{k+s}} = E\varepsilon_{S_{k+s}}\varepsilon'_{S_{k+s}} = \mathbf{D}_{S_{k+s}}\mathbf{P}_{S_{k+s}}\mathbf{D}_{S_{k+s}} \quad (5)$$

see Bollerslev (1990), where

- ▶  $\mathbf{D}_{S_{k+s}}$  is a diagonal matrix of standard deviations and
- ▶  $\mathbf{P}_{S_{k+s}}$  is a positive definite correlation matrix.

$\mathbf{D}_{S_{k+s}}$  and  $\mathbf{P}_{S_{k+s}}$  may vary both with  $s$  and  $k$ .

# The VSSMC-AR-X model

## The covariance equation

Specifically,  $\mathbf{D}_{S_{k+s}} = \text{diag}(\sigma_{s1}(u_{ks}), \dots, \sigma_{sN}(u_{ks}))$ , where

$$\sigma_{sl}^2(u_{ks}) = \sigma_{sl0}^2 + \sum_{i=1}^{r_{sl}} \omega_{sli} g_{sli}(u_{ks}; \gamma_{sli}^{(v)}, c_{sli}^{(v)}) \quad (6)$$

for  $l = 1, \dots, N$ . In (6),

$$g_{sli}(u_{ks}; \gamma_{sli}^{(v)}, c_{sli}^{(v)}) = (1 + \exp\{-\gamma_{sli}^{(v)}(u_{ks} - c_{sli}^{(v)})\})^{-1} \quad (7)$$

or

$$g_{sli}(u_{ks}; \gamma_{sli}^{(v)}, c_{sli}^{(v)}) = 1 - \exp\{-\gamma_{sli}^{(v)}(u_{ks} - c_{sli}^{(v)})^2\}. \quad (8)$$

In (7) and (8),  $\gamma_{sli}^{(v)} > 0$ ,  $i = 1, \dots, r_{sl}$ ;  $s = 1, \dots, S$ .

# The VSSMC-AR-X model

The covariance equation

Positivity condition for (6):

$$\sigma_{s|0}^2 + \sum_{i=1}^{r_{sl}} \omega_{sli} g_{sli}(r; \gamma_{sli}^{(v)}, c_{sli}^{(v)}) > 0$$

for  $\forall r \in [0, 1]$ ,  $s$  and  $k$ .



# The VSSMC-AR-X model

## The covariance equation

The time-varying error correlation matrix  $\mathbf{P}_{Sk+s}$  in (5) for season  $s$  has the following form, see, for example, He et al. (2021),

$$\mathbf{P}_{Sk+s} = \{1 - g_s^{(\rho)}(u_{ks})\} \mathbf{P}_{(s1)} + g_s^{(\rho)}(u_{ks}) \mathbf{P}_{(s2)} \quad (9)$$

where  $\mathbf{P}_{(s1)}$  and  $\mathbf{P}_{(s2)}$  are  $N \times N$  positive definite correlation matrices for season  $s$ .

# The VSSMC-AR-X model

## The covariance equation

In (9),

$$g_s^{(\rho)}(u_{ks}; \gamma_s^{(\rho)}, c_s^{(\rho)}) = (1 + \exp\{-\gamma_s^{(\rho)}(u_{ks} - c_s^{(\rho)})\})^{-1} \quad (10)$$

or

$$g_s^{(\rho)}(u_{ks}; \gamma_s^{(\rho)}, c_s^{(\rho)}) = 1 - \exp\{-\gamma_s^{(\rho)}(u_{ks} - c_s^{(\rho)})^2\}. \quad (11)$$

- ▶ Since both (10) and (11) are bounded between zero and one, as a convex combination of two positive definite correlation matrices  $\mathbf{P}_{Sk+s}$  is positive definite for all  $Sk + s$ .

# The VSSMC-AR-X model

## The covariance equation

- ▶ NOTE: In the application,  $\mathbf{P}_{S_{k+s}} = \mathbf{P}_s$ .
- ▶ (Another simplification:  $\Phi_h$ ,  $h = 1, \dots, p$ , in the mean equation are diagonal matrices.)

# The VSSMC-AR-X model

## Testing linearity of the mean equation

Before fitting the VSSMC-AR-X model to the data, it is necessary to test linearity.

- ▶ From the point of view of **the application**, one has to know whether or not the temperature series are nonstationary.
- ▶ From the **statistical point of view**, testing is necessary because the  $l$ th equation is not identified if the linearity hypothesis  $\delta_{sl}(u_{ks}) = \delta_{s/0}$  in (2) holds for any  $s = 1, \dots, S$ .

Example:

$$\delta_{sl}(u_{ks}) = \delta_{s/0} + \delta_{s/1}(1 + \exp\{-\gamma_{s/1}(u_{ks} - c_{s/1})\})^{-1}.$$

# The VSSMC-AR-X model

## Testing linearity of the mean equation

- ▶ The test is based on approximating the alternative with a Taylor expansion around the null hypothesis  $\gamma_{sli} = 0$ , see Luukkonen et al. (1988).
- ▶ That is, testing zero transitions against one let  $H_0: \gamma_{s/1} = 0$  and use the approximation (in this example a third-order one)

$$\delta_{sl}(u_{ks}) \approx \delta_{s/0}^* + \delta_{s/1}^* u_{ks} + \delta_{s/2}^* u_{ks}^2 + \delta_{s/3}^* u_{ks}^3.$$

This results in a new null hypothesis  $H'_0$ :

$\delta_{s/1}^* = \delta_{s/2}^* = \delta_{s/3}^* = 0$ . Can be tested by a standard  $\chi^2$ -test.

# The VSSMC-AR-X model

## Testing linearity of the mean equation

Testing is carried out in stages as in He et al. (in press).

- ▶ First test the null hypothesis against one transition, i.e.,  $q_{sI} = 1$  in (2). Do this separately for  $s = 1, \dots, S$ .
- ▶ If the null hypothesis is rejected for at least one  $s$ , estimate the equation with one transition and for the seasons with one transition, test against two.
- ▶ Proceed until the first non-rejection.

# The VSSMC-AR-X model

## ML estimation of parameters

- ▶ Estimate the mean equation first, then the variances and correlations.
- ▶ The variances and correlations are estimated **jointly** (no two-stage estimation).
  - ▶ Constancy of variances has to be tested before estimating the variance equations (identification problem).
- ▶ Asymptotic properties of ML estimators, see He et al. (in press).

# The VSSMC-AR-X model

Joint estimation of variance and correlation parameters

Notation:

- ▶  $\boldsymbol{\theta}_s^{(v)} = (\boldsymbol{\theta}_{1s}^{(v)'} , \dots , \boldsymbol{\theta}_{Ns}^{(v)'})'$ ,  $s = 1, \dots, S$ , is the vector of variance parameters for season  $s$ ,  $\boldsymbol{\theta}_{ns}^{(v)}$ ,  $n = 1, \dots, N$ , is the vector of parameters for season  $s$  in the  $n$ th equation
- ▶  $\boldsymbol{\theta}_s^{(c)}$  is the vector of parameters in the correlation matrix  $\mathbf{P}_{Sk+s} = \mathbf{P}_{Sk+s}(\boldsymbol{\theta}_s^{(c)})$ .



# The VSSMC-AR-X model

## Joint estimation of variance and correlation parameters

For season  $s$ , estimation proceeds as follows (Silvennoinen and Teräsvirta, in press).

1. Estimate the parameters in  $\boldsymbol{\theta}_s^{(v)} = (\boldsymbol{\theta}_{1s}^{(v)'} , \dots , \boldsymbol{\theta}_{Ns}^{(v)'})'$ ,  $s = 1, \dots, S$ , equation by equation, assuming  $\mathbf{P}_{S_{k+s}}(\boldsymbol{\theta}_s^{(c)}) = \mathbf{I}_N$ . Denote the estimate  $\hat{\boldsymbol{\theta}}_s^{(v,1)} = (\hat{\boldsymbol{\theta}}_{1s}^{(v,1)'}, \dots, \hat{\boldsymbol{\theta}}_{Ns}^{(v,1)'})'$ . This means that the deterministic components  $\sigma_{ns}^2(u_{ks})$ ,  $n = 1, \dots, N$ , have been estimated once.
2. Estimate  $\mathbf{P}_{S_{k+s}}(\boldsymbol{\theta}_s^{(c)})$ , given  $\boldsymbol{\theta}_s^{(v)} = \hat{\boldsymbol{\theta}}_s^{(v,1)}$ ,  $s = 1, \dots, S$ . This requires a separate iteration because  $\mathbf{P}_{S_{k+s}}(\boldsymbol{\theta}_s^{(c)})$  is nonlinear in parameters, see (9), and (10) or (11). Denote the estimate  $\mathbf{P}_{S_{k+s}}(\hat{\boldsymbol{\theta}}_s^{(c,1)})$ .

# The VSSMC-AR-X model

Joint estimation of variance and correlation parameters

3. Re-estimate  $\theta_s^{(v)}$  assuming  $\mathbf{P}_{S_{k+s}}(\theta_s^{(c)}) = \mathbf{P}_{S_{k+s}}(\hat{\theta}_s^{(c,1)})$ . This yields  $\theta_s^{(v)} = \hat{\theta}_s^{(v,2)}$ . Then re-estimate  $\mathbf{P}_{S_{k+s}}(\theta_s^{(c)})$  given  $\theta_s^{(v)} = \hat{\theta}_s^{(v,2)}$ . Iterate until convergence. Let the result after  $R_1$  iterations be  $\theta_s^{(v)} = \hat{\theta}_s^{(v,R_1)}$  and  $\mathbf{P}_{S_{k+s}}(\theta_s^{(c)}) = \mathbf{P}_{S_{k+s}}(\hat{\theta}_s^{(c,R_1)})$ .

This completes the first iteration of the algorithm.

4. Re-estimate first the mean parameters given  $\hat{\theta}_s^{(v,R_1)}$  and  $\mathbf{P}_{S_{k+s}}(\hat{\theta}_s^{(c,R_1)})$ ,  $s = 1, \dots, S$ , then jointly variance and correlation parameters using  $\theta_s^{(v)} = \hat{\theta}_s^{(v,R_1)}$  and  $\mathbf{P}_{S_{k+s}}(\theta_s^{(c)}) = \mathbf{P}_{S_{k+s}}(\hat{\theta}_s^{(c,R_1)})$  as starting values for the latter. This completes the second iteration of the algorithm.

5. Iterate until convergence (means, variances, correlations).

# The data

28 weather stations in Europe, monthly time series for the years 1823–2015

- ▶ from Paris in the west to Kazan in the east,
- ▶ from Arkhangelsk in the north to Milan in the south,
- ▶ stations with data available from 1823 but with rather large numbers of missing variables were not included.

# The data

28 weather stations in Europe, monthly time series for the years 1823–2015

- ▶ from Paris in the west to Kazan in the east,
- ▶ from Arkhangelsk in the north to Milan in the south,
- ▶ stations with data available from 1823 but with rather large numbers of missing variables were not included.
- ▶ (Tallinn not included: the series begins 1828 with gaps 1876–1880 and 1916–1920.)

# The data

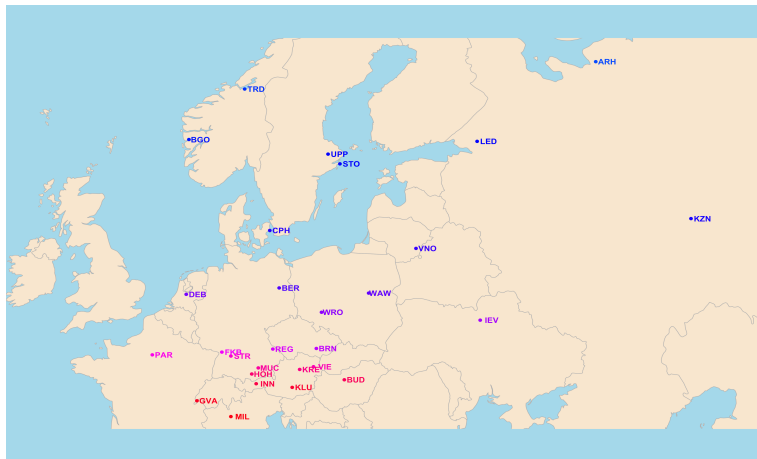


Figure: Map showing the locations of the 28 cities and towns from Arkhangelsk in the north to Milan in the south

# Densities of temperature series

Trondheim, Uppsala, Stockholm, Copenhagen, Vilnius, Berlin, Warsaw and De Bilt

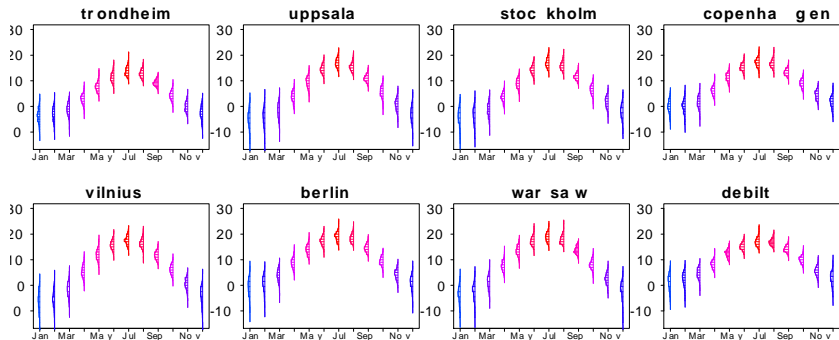


Figure: Densities (histograms) by month of eight temperature series, ordered by latitude, from Trondheim to De Bilt. Source: He et al. (in press)

## The data

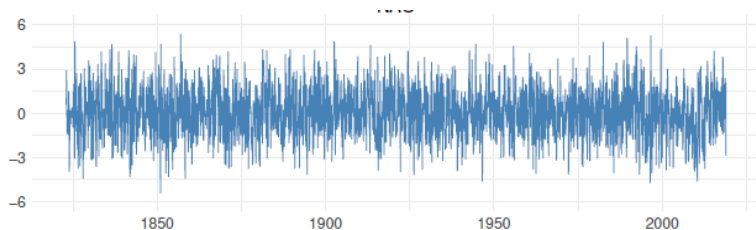


Figure: The monthly North Atlantic Oscillation index 1823–2015.  
Source: Jones, Jónsson and Wheeler (1997), extended to 2015.

## Results in the literature

$y_t$  is nonstationary and dynamic

$x_t$  is stationary (and dynamic)

- ▶ Hurrell and van Loon (1997), Hurrell (1996):  
Estimate  $\beta_1$  from  $y_t = \beta_0 + \beta_1 x_t + e_t$ , where  $t$  is the aggregate for winter months
- ▶ Osborn (2011):  
Estimate  $\beta_1$  from  $y_t = \beta_0 + \beta_1 x_t + e_t$  by month (a separate equation for each month)
- ▶ Iles and Hegerl (2017):  
Estimate  $\delta_1$  from  $x_t = \delta_0 + \delta_1 y_t + e'_t$  (they have a dense grid; they probably mean regressing  $y$  on  $x$ )



# Results from the VSSMC-AR model (He et al., in press)

## Shifting means by month, North

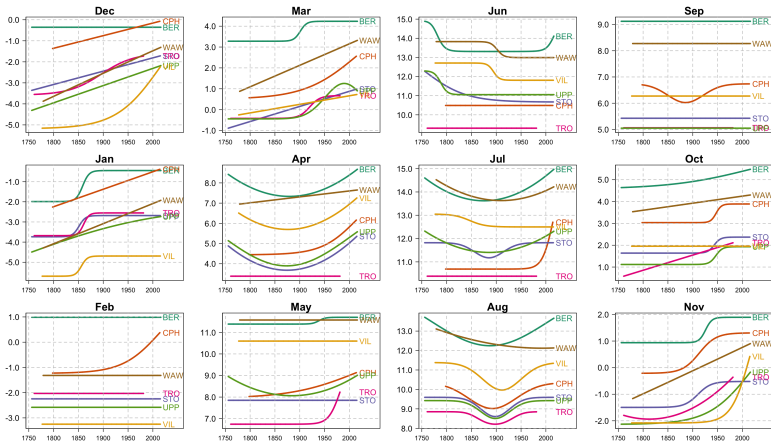


Figure: Estimated monthly temperature shifts for the locations in North. From north to south: Trondheim (TRO), Uppsala (UPP), Stockholm (STO), Copenhagen (CPH), Vilnius (VIL), Berlin (BER) and Warsaw (WAW).

# Results from the VSSMC-AR model (He et al., in press)

## Shifting means by month, West

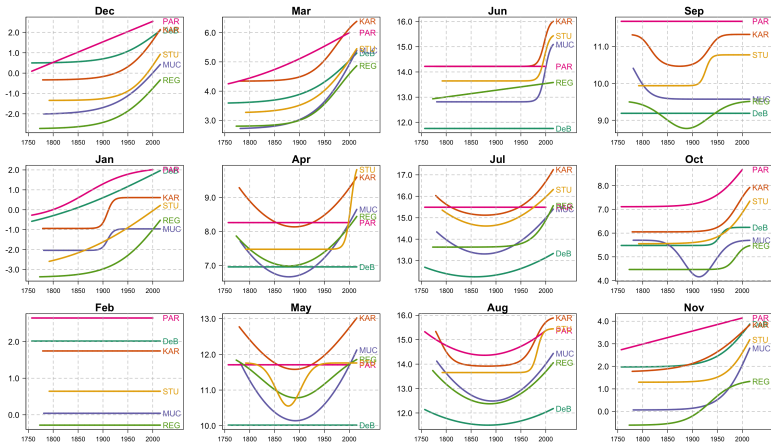


Figure: Estimated monthly temperature shifts for the locations in West. From west to east: Paris (PAR), De Bilt (DeB), Karlsruhe (KAR), Stuttgart (STU), Munich (MUN) and Regensburg (REG).

# Results from the VSSMC-AR model (He et al., in press)

## Shifting means by month, East-South

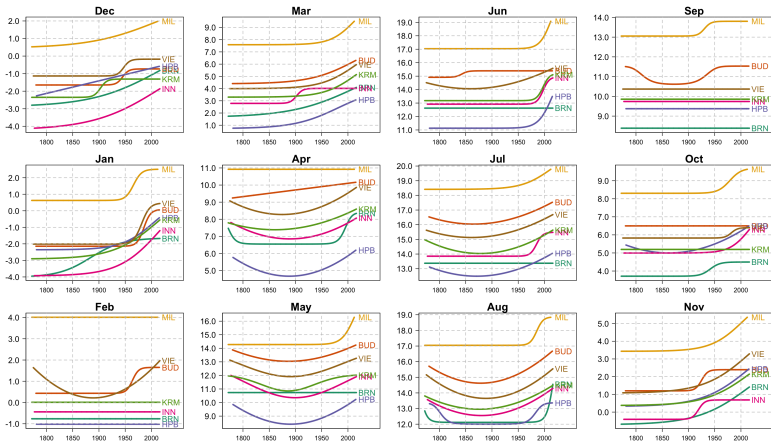
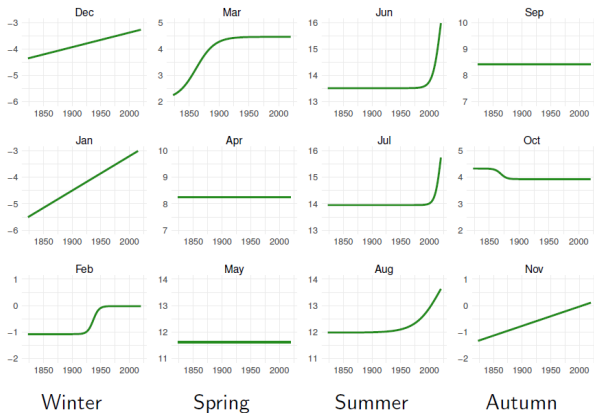


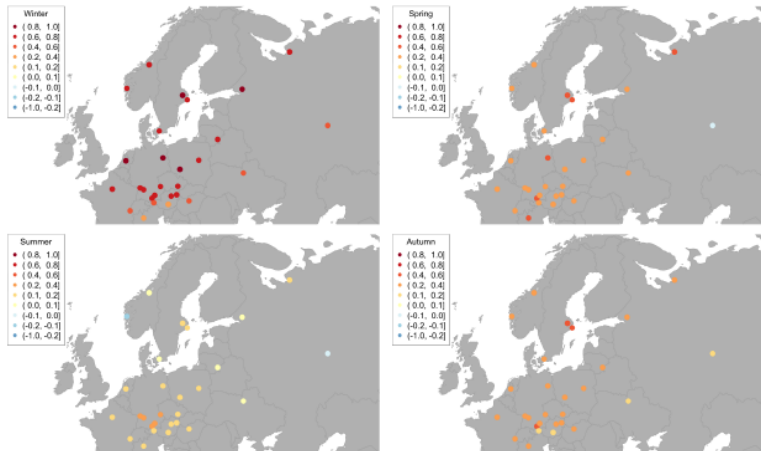
Figure: Estimated monthly temperature shifts for the locations in East/South. From west to east: Milan (MIL), Hohenpeissenberg (HPB), Innsbruck (INN), Kremsmünster (KRE), Vienna (VIE), Brno (BRN) and Budapest (BUD).

# Results from the VSSMC-AR-X model

The shifting mean: Klagenfurt

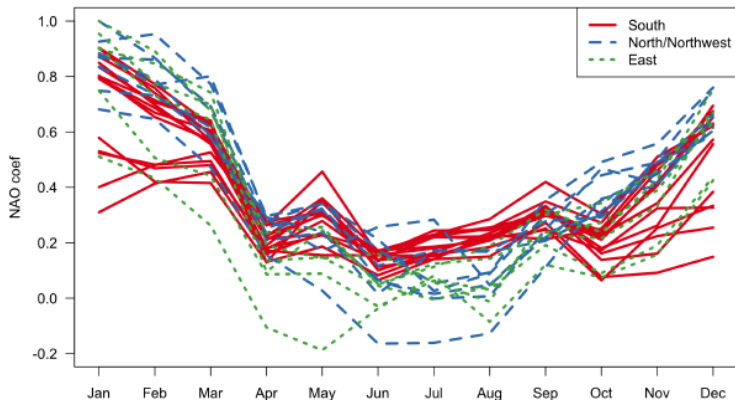


# Results from the VSSMC-AR-X model



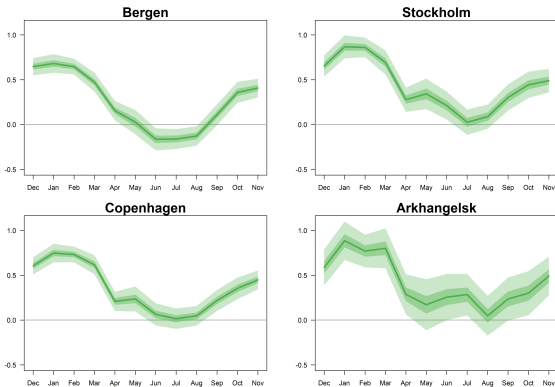
Seasonal averages of coefficient estimates of  $x_t$ . Top left: Winter (Dec-Feb), Top right: Spring (Mar-May), Bottom left: Summer (Jun-Aug), Bottom right: Autumn (Sep-Nov)

## Results from the VSSMC-AR-X model



Estimates of the coefficient of  $x_t$  over the year in (1): Dashed lines: North/Northwest (Arkhangelsk, Bergen, Trondheim, Uppsala, Stockholm, Copenhagen, De Bilt, Berlin), Dotted lines: East (Kyiv, St Petersburg, Wroclaw, Warsaw, Vilnius, Kazan), Solid lines: South (the remaining 14 stations)

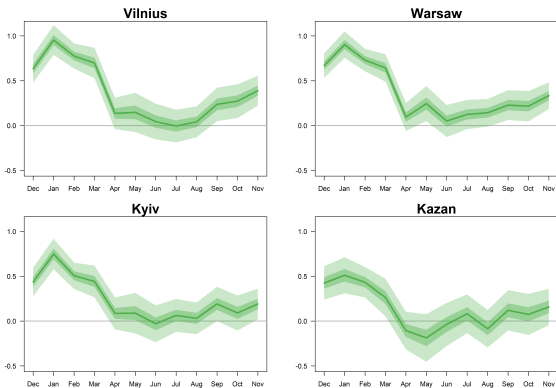
# Results from the VSSMC-AR-X model



Estimates of the coefficient of  $x_t$  over the year in (1), solid line; 50% confidence level, dark green; 95% confidence level, light green.

North/Northwest: Top panel: Bergen and Stockholm, Bottom panel: Copenhagen and Arkhangelsk

# Results from the VSSMC-AR-X model

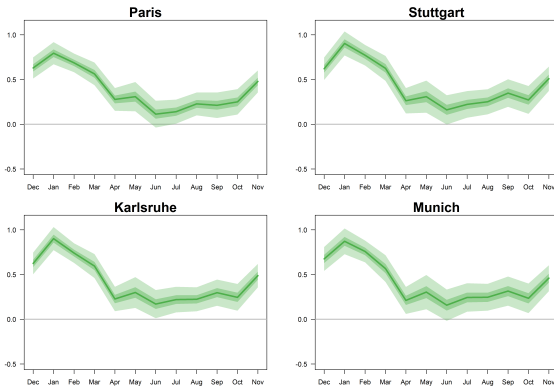


Estimates of the coefficient of  $x_t$  over the year in (1), solid line, 50% confidence level dark green, 95% confidence level, light green.

North/Northeast: Top panel: Vilnius and Warsaw, Bottom panel: Kyiv and Kazan

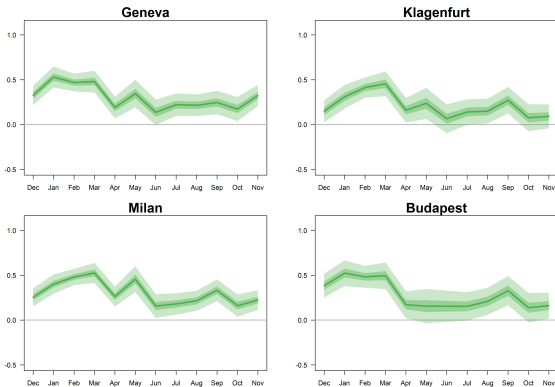


# Results from the VSSMC-AR-X model



Estimates of the coefficient of  $x_t$  over the year in (1), solid line; 50% confidence level, dark green; 95% confidence level, light green. West: Top panel: Paris and Stuttgart, Bottom panel: Karlsruhe and Munich

# Results from the VSSMC-AR-X model



Estimates of the coefficient of  $x_t$  over the year in (1), solid line; 50% confidence level, dark green; 95% confidence level, light green.

South/Southeast: Top panel: Geneva and Klagenfurt, Bottom panel: Milan and Budapest

# Results from the VSSMC-AR-X model

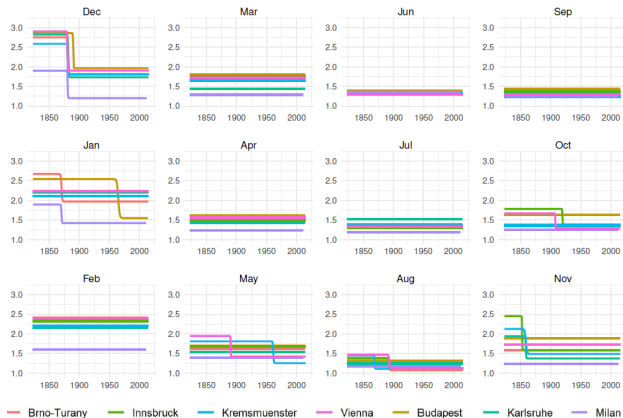
## Time-varying standard deviations



Standard deviation estimates: North (Arkhangelsk, Berlin, De Bilt, Trondheim, Copenhagen, Stockholm, Uppsala)

# Results from the VSSMC-AR-X model

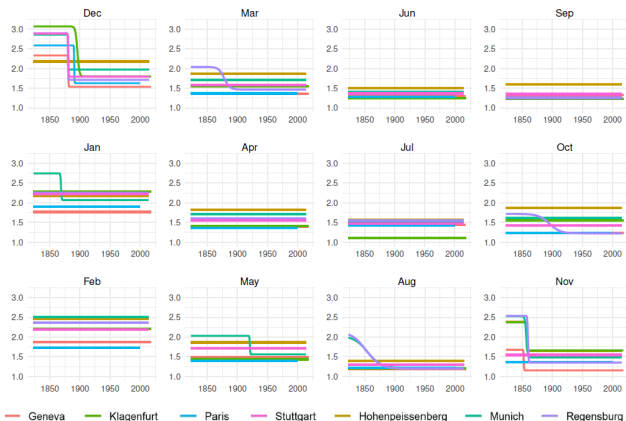
## Time-varying standard deviations



Standard deviation estimates: Southeast (Brno-Turany, Innsbruck, Vienna, Budapest, Karlsruhe, Milan)

# Results from the VSSMC-AR-X model

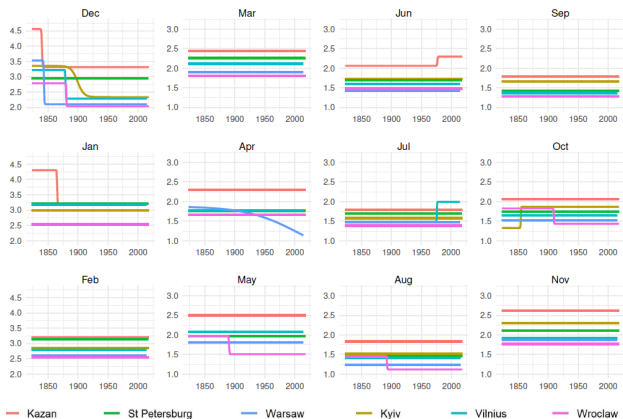
## Time-varying standard deviations



Standard deviation estimates: Southwest (Geneve, Klagenfurt, Paris, Stuttgart, Hohenpeissenberg, Munich, Regensburg)

# Results from the VSSMC-AR-X model

## Time-varying standard deviations



Standard deviation estimates: East (Kazan, St Petersburg, Warsaw, Kyiv, Vilnius, Wroclaw)

## Error correlations: Spatial relationships

Generalise Haslett and Raftery (1989):

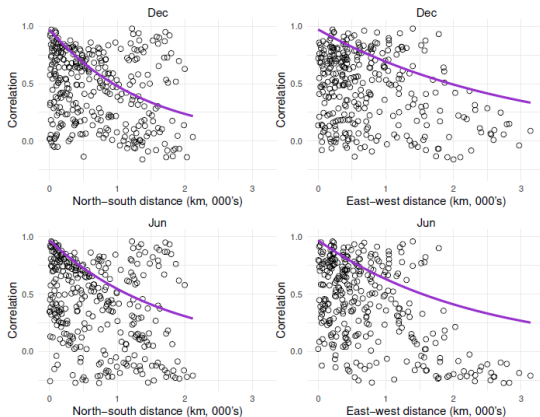
- ▶  $d_{ij}^{NS}$  = latitudinal (north-south) distance between stations  $i$  and  $j$
- ▶  $d_{ij}^{EW}$  = longitudinal (east-west) distance between stations  $i$  and  $j$
- ▶  $d_{ij}^H$  = absolute difference in elevation between stations  $i$  and  $j$
- ▶  $\hat{\rho}_{sij}$  = residual correlation between stations  $i$  and  $j$  for month  $s$

Minimise

$$Q_s^{(c)} = \min_{\theta_s^{(c)}} (\hat{\rho}_{sij} - \alpha_s \exp\{-(\beta_s^{NS} d_{ij}^{NS} + \beta_s^{EW} d_{ij}^{EW} + \beta_s^H d_{ij}^H)\})^2$$

where  $\alpha_s$  is the 'nugget effect' (Haslett and Raftery, 1989)

# Error correlations: Spatial relationships



Estimated correlations and residual correlations against distance (in 1000K). (i) December: Top left: north-south, top right: east-west, (ii) June: Bottom left: north-south, bottom right: east-west



# Error correlations: Spatial relationships

## Correlations vs. absolute differences in elevation

- ▶ Statistically significant (but weak) relationship during the extended winter (Nov-Mar).
- ▶ No relationship for the remaining months.

# Conclusions

- ▶ The NAO affects the weather in Europe:
  - ▶ The effect is strongest in the winter and declines (not monotonically) towards the summer.
  - ▶ The winter effect is stronger in northern latitudes than in the south.
  - ▶ In the summer the effect is still significantly different from zero in the west, not so in the east.
- ▶ The estimated model makes it possible to
  - ▶ forecast temperatures given forecasts on the NAO

# Conclusions

- ▶ The NAO affects the weather in Europe:
  - ▶ The effect is strongest in the winter and declines (not monotonically) towards the summer.
  - ▶ The winter effect is stronger in northern latitudes than in the south.
  - ▶ In the summer the effect is still significantly different from zero in the west, not so in the east.
- ▶ The estimated model makes it possible to
  - ▶ forecast temperatures given forecasts on the NAO (Warning!)

# Conclusions

- ▶ The NAO affects the weather in Europe:
  - ▶ The effect is strongest in the winter and declines (not monotonically) towards the summer.
  - ▶ The winter effect is stronger in northern latitudes than in the south.
  - ▶ In the summer the effect is still significantly different from zero in the west, not so in the east.
- ▶ The estimated model makes it possible to
  - ▶ forecast temperatures given forecasts on the NAO (**Warning!**)
  - ▶ run counterfactuals (e.g., what happens to temperatures if instead of 'high' NAO there is 'low' NAO?)

# Conclusions

- ▶ The NAO affects the weather in Europe:
  - ▶ The effect is strongest in the winter and declines (not monotonically) towards the summer.
  - ▶ The winter effect is stronger in northern latitudes than in the south.
  - ▶ In the summer the effect is still significantly different from zero in the west, not so in the east.
- ▶ The estimated model makes it possible to
  - ▶ forecast temperatures given forecasts on the NAO (**Warning!**)
  - ▶ run counterfactuals (e.g., what happens to temperatures if instead of 'high' NAO there is 'low' NAO?)
- ▶ Distances between pairs of weather stations have a weak effect on error correlations.

## Future work

- ▶ Climatologists know that the NAO also affects precipitation, at least in the winter.
- ▶ The VSSMC-AR-X model may be applied to investigating
  - ▶ seasonality, nonlinearity and nonstationarity of long monthly European precipitation series,
  - ▶ the effects of the NAO on precipitation using such series.
- ▶ This investigation has just begun.